

Introduction

Three regions at stake to understand auroral process:

- **Ionosphere**: ionized plasma in the high atmosphere (80 km to 400 km above Earth's surface)
- **Acceleration region**: particles are expected to gain energy (several 1,000 km above Earth's surface)
- **Generator**: conversion of mechanical energy to electromagnetic energy (around 8 Earth's radii).

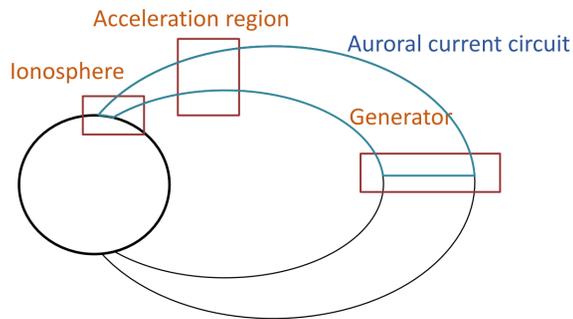


Figure: The three key regions in the auroral current circuit

Objectives

Goal: study of the magnetosphere-ionosphere coupling via field-aligned currents (FAC) carried by shear Alfvén.

Hypothesis:

- The acceleration region is inside the ionosphere
- The geomagnetic field lines are supposed to be straight lines with a dip angle I such that in a non-tilted dipole approximation

$$\sin I = \frac{2 \cos \theta}{\sqrt{1 + 3(\cos \theta)^2}}$$

with θ the colatitude.

Notations:

- \vec{E}_\perp and j_\parallel the electric field orthogonal and the current parallel to the magnetic field lines
- $\Sigma_\parallel, \Sigma_P, \Sigma_H$ parallel, Pederson and Hall height-integrated conductivities
- Σ_A Alfvén conductance ($1/\mu_0 V_A$)
- K a constant such that $K \sim 10^{-10} \Omega^{-1} \cdot m^{-2}$.

Model & Methods

Model for the reflection at the ionosphere

The reflection at the ionosphere is governed by:

$$\vec{\nabla}_\perp \cdot (\Sigma^+ \cdot \vec{E}_\perp^+ + \Sigma^- \cdot \vec{E}_\perp^- - \Sigma \cdot \vec{\nabla}_\perp \left[-\frac{j_\parallel}{K} \right]) = 0$$

with

$$\Sigma^\pm = \begin{pmatrix} \Sigma_{\theta\theta} \pm \Sigma_A & \Sigma_{\theta\varphi} \\ \Sigma_{\theta\varphi} & \Sigma_{\varphi\varphi} \pm \Sigma_A \end{pmatrix}, \quad \Sigma = \begin{pmatrix} \Sigma_{\theta\theta} & \Sigma_{\theta\varphi} \\ \Sigma_{\theta\varphi} & \Sigma_{\varphi\varphi} \end{pmatrix}$$

where

$$\begin{cases} \Sigma_{\theta\theta} = \frac{\Sigma_P \Sigma_\parallel}{\Sigma_P + (\Sigma_\parallel - \Sigma_P) (\sin I)^2} \\ \Sigma_{\theta\varphi} = \frac{\Sigma_P \Sigma_\parallel \sin I}{\Sigma_P + (\Sigma_\parallel - \Sigma_P) (\sin I)^2} \\ \Sigma_{\varphi\varphi} = \Sigma_P + \frac{\Sigma_\parallel^2 (1 - (\sin I)^2)}{\Sigma_P + (\Sigma_\parallel - \Sigma_P) (\sin I)^2} \end{cases}$$

Developing leads to an equation of type (1)

$$A(\theta, \varphi) \frac{\partial^2 \phi}{\partial \theta^2} + B(\theta, \varphi) \frac{\partial \phi}{\partial \theta} + C(\theta, \varphi) \frac{\partial^2 \phi}{\partial \varphi^2} + D(\theta, \varphi) \frac{\partial \phi}{\partial \varphi} = F(\theta, \varphi)$$

in spherical coordinates.

Results

Evolution of field-aligned currents

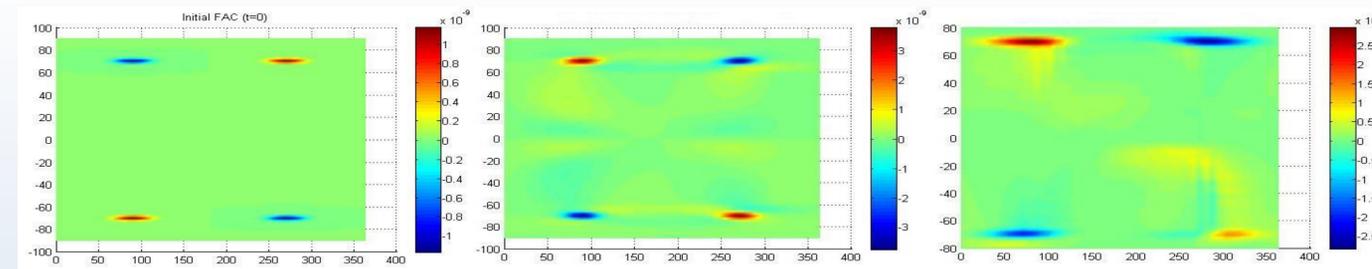


Figure: Initial FAC (left) and after MI-coupling with theoretical (middle) and IRI (right) conductivities

Potential in the ionosphere

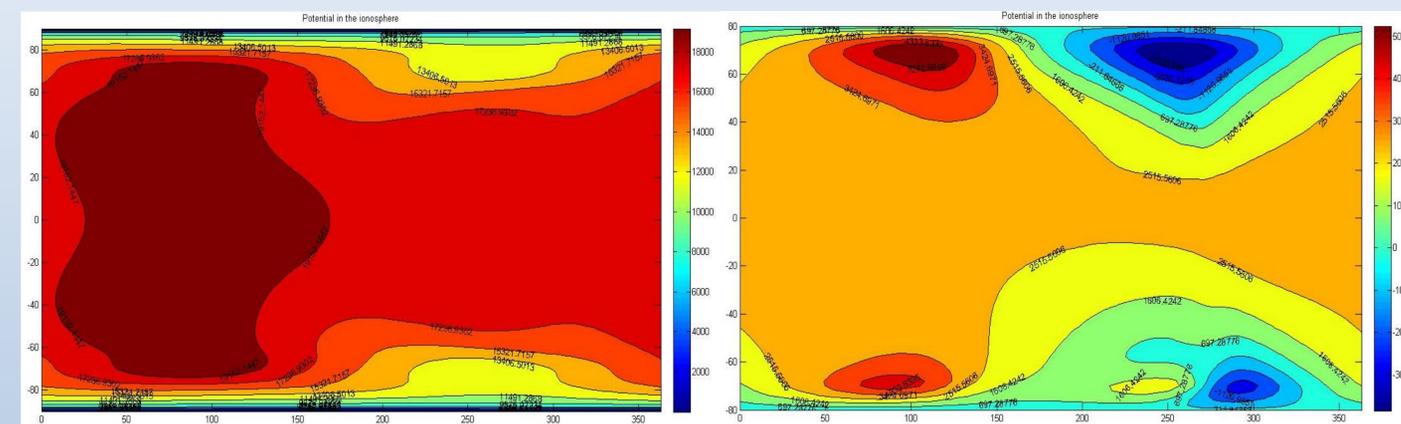


Figure: Potential in the ionosphere with theoretical (left) and IRI \bar{e} conductivities (right)

Resolution

The equations of type (1) are solved with *finite differences* and the time-dependent equation with a *forward integration in time*.

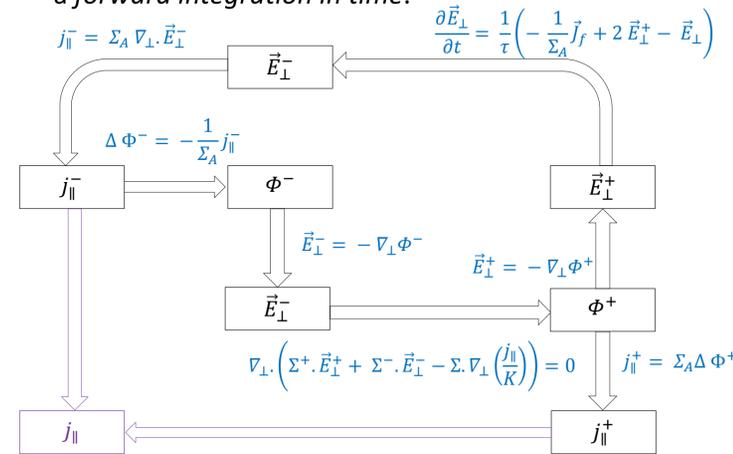


Figure: Summary of the equations to solve in the MI-coupling. The process starts with a given j_\parallel^- and ends when it has converged

Height-integrated conductivities

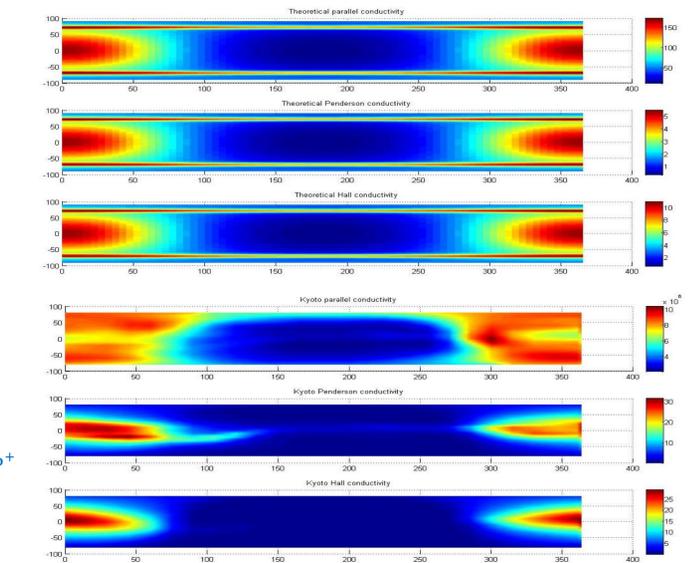


Figure: Height-integrated conductivities profiles: theoretical (top) and based on International Reference Ionosphere (bottom)

Conclusion

Two major differences:

- Field-aligned currents reach lower latitudes with IRI conductivities
- The potential is greater and iso-potential lines reach lower latitudes with theoretical conductivities.

Limitations:

- Key regions are supposed to be thin sheets parallel
- Field lines are modelled by straight lines
- Fluid description not always valid

Acknowledgements

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References

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